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2 PROF. OWEN K ATKIN (Orcid ID : 0000-0003-1041-5202)

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8 **The validity of optimal leaf traits modelled on environmental conditions**

9 Keith J. Bloomfield*¹, I. Colin Prentice^{2,3}, Lucas A. Cernusak⁴, Derek Eamus⁵, Belinda E.
10 Medlyn⁶, Rizwana Rumman⁵, Ian J. Wright², Matthias M. Boer⁶, Peter Cale⁷, James
11 Cleverly^{5,8}, John J.G. Egerton¹, David S. Ellsworth⁶, Bradley J. Evans⁹, Lucy S. Hayes¹,
12 Michael F. Hutchinson¹⁰, Michael J. Liddell¹¹, Craig Macfarlane¹², Wayne S. Meyer¹³,
13 Henrique F. Togashi², Tim Wardlaw¹⁴, Lingling Zhu^{1,15} and Owen K. Atkin^{1,15}

14

15 ¹Division of Plant Sciences, Research School of Biology, Building 46, The Australian National
16 University, Canberra, ACT 2601, Australia; ²Department of Biological Sciences, Macquarie
17 University, North Ryde, NSW 2109, Australia; ³AXA Chair of Biosphere and Climate Impacts,
18 Grand Challenges in Ecosystems and the Environment and Grantham Institute—Climate
19 Change and the Environment, Department of Life Sciences, Imperial College London, Silwood
20 Park Campus, Buckhurst Road, Ascot SL5 7PY, UK; ⁴Department of Marine and Tropical
21 Biology, James Cook University, Cairns, Qld 4878, Australia; ⁵School of Life Sciences,
22 University of Technology Sydney, NSW 2007, Australia; ⁶Hawkesbury Institute for the
23 Environment, Western Sydney University, Penrith, NSW 2751, Australia; ⁷Australian Landscape
24 Trust, Renmark, SA 5341, Australia; ⁸TERN (Terrestrial Ecosystem Research Network,
25 University of Technology Sydney); ⁹Faculty of Agriculture and Environment, Department of
26 Environmental Sciences, University of Sydney, Sydney, NSW 2006, Australia; ¹⁰Fenner School

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27 of Environment and Society, Australian National University, Canberra, ACT 2601, Australia;
28 ¹¹Centre for Tropical, Environmental, and Sustainability Sciences, James Cook University,
29 Cairns, Qld, Australia; ¹²CSIRO Land and Water, Private Bag 5, Wembley, WA 6913, Australia;
30 ¹³Earth and Environmental Sciences, University of Adelaide, Adelaide, SA 5064, Australia;
31 ¹⁴ARC Centre for Forest Value, University of Tasmania, Hobart, TAS 7005, Australia; ¹⁵ARC
32 Centre of Excellence in Plant Energy Biology, Research School of Biology, Building 134, The
33 Australian National University, Canberra, ACT 2601, Australia.

34 * Corresponding author: Keith.Bloomfield@anu.edu.au

35 Tel: +44 28 71341685

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38

39 Lucas Cernusak 0000-0002-7575-5526

40 James Cleverly 0000-0002-2731-7150

41 Derek Eamus 0000-0003-2765-8040

42 Ian Wright 0000-0001-8338-9143

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44 **Summary**

45

- 46 • The ratio of leaf intercellular to ambient CO₂ (γ) is modulated by stomatal conductance
47 (g_s). These quantities link carbon (C) assimilation with transpiration, and along with
48 photosynthetic capacities (V_{cmax} and J_{max}) are required to model terrestrial C uptake. We
49 use optimisation criteria based on the growth environment to generate predicted values
50 of photosynthetic and water-use efficiency traits and test these against a unique dataset.
- 51 • Leaf gas-exchange parameters and carbon isotope discrimination were analysed in
52 relation to local climate across a continental network of study sites. Sun-exposed leaves
53 of 50 species at seven sites were measured in contrasting seasons.

- 54 • Values of χ predicted from growth temperature and vapour pressure deficit were closely
55 correlated to ratios derived from C isotope ($\delta^{13}\text{C}$) measurements. Correlations were
56 stronger in the growing season. Predicted values of photosynthetic traits, including
57 carboxylation capacity (V_{cmax}), derived from $\delta^{13}\text{C}$, growth temperature and solar
58 radiation, showed meaningful agreement with inferred values derived from gas-exchange
59 measurements. Between-site differences in water use efficiency were, however, only
60 weakly linked to the plant's growth environment and did not show seasonal variation.
- 61 • These results support the general hypothesis that many key parameters required by Earth
62 system models are adaptive and predictable from plants' growth environments.
- 63
- 64

65 **Keywords:**

66 aridity, photosynthesis, stable isotopes, stomatal conductance (g_s), temperature, water-use
67 efficiency.

68

69 **Introduction**

70 All healthy leaves ultimately fulfil the same functions of light capture and gas exchange. In C_3
71 plants, the ratio (χ) of leaf intercellular (C_i) to ambient (C_a) partial pressure of CO_2 depends on a
72 balance between stomatal conductance (g_s) and photosynthesis, which in turn depends on C_i , on
73 leaf-level capacities for carboxylation (V_{cmax}) and electron transport (J_{max}), and on temperature-
74 dependent kinetic constants. We can have confidence in our ability to measure and monitor C_a ,
75 but knowledge of associated C_i - made possible by reliable estimates of χ - is crucial to
76 understanding the leaf's carbon economy. There are large variations in χ , g_s , V_{cmax} and J_{max}
77 across species and environments (Schulze *et al.*, 1994; De Kauwe *et al.*, 2016; Wang *et al.*, 2017b).
78 Faced with huge global diversity in plant species (and the morphology, physiology and longevity
79 of leaves) there is great benefit in quantifying unifying trends that may be widely applied in
80 global-scale models. The worldwide leaf economics spectrum (LES) is a notable example
81 (Wright *et al.*, 2004), most importantly describing a trade-off between specific leaf area (the ratio
82 of leaf area to dry mass) and leaf longevity. Considerable variation in the gas exchange
83 parameters listed above is, however, unrelated to the LES. Given the fundamental importance
84 of these parameters for modelling land-atmosphere carbon and water exchange, it is vital that
85 tools be developed to predict variations in χ , g_s , V_{cmax} and J_{max} .

86 Current models of vegetation and land-surface processes, including those embedded in Earth
87 system models (ESMs), typically prescribe values of V_{cmax} and J_{max} for each plant functional type
88 (PFT) and use empirical models to predict χ and g_s . This approach has major limitations because
89 (a) there is at least as much variation in V_{cmax} and J_{max} within as among PFTs (e.g. Verheijen *et al.*,
90 2013; Rogers, 2014) and (b) the empirical models require the specification of further parameters
91 for PFTs. An alternative approach is to seek general functional relationships that allow these
92 photosynthetic parameters to be predicted from environmental information (Reich *et al.*, 2007;
93 Ali *et al.*, 2015). A promising route to defining such relationships is provided by optimality
94 hypotheses, couched in quantitative terms that allow testing against field measurements (Mäkelä
95 *et al.*, 2002; Prentice *et al.*, 2014; Prentice *et al.*, 2015). There is a need to test the predictions
96 emanating from such models against empirical data drawn from as wide as possible a range of
97 natural settings.

98 To balance water loss against carbon gain, the plant must control g_s which modulates both rates.
99 When stomata are open, the leaf-to-air vapour pressure deficit (D) provides the gradient that
100 drives transpiration (E). Ratios of carbon assimilation (A) to E , provide one measure of the
101 plant's water use efficiency (WUE) and for C_3 plants vary from 2 to 11 mmol C mol⁻¹ H₂O
102 (Lambers *et al.*, 2008), illustrating the high water cost of C gain. Cowan and Farquhar (1977)
103 introduced an optimality hypothesis for stomatal regulation, defining a parameter λ that may be
104 thought of as an exchange rate between water and C. Their hypothesis proposes that, over the
105 course of a single day, optimal stomatal behaviour minimises E for a given A such that the
106 relative sensitivities of E and A to changes in g_s remain constant:

$$\frac{\partial E / \partial g_s}{\partial A / \partial g_s} = \frac{\partial E}{\partial A} = \lambda \quad (1)$$

107 Cowan and Farquhar pointed out that this exchange rate must vary with ontogeny and changing
108 environment – no single value could be appropriate at all life stages and under all conditions.
109 Subsequent controlled-environment studies across a range of tree species have confirmed that
110 estimated values of λ , derived from empirical models, vary with leaf temperature (T), D ,
111 irradiance, soil water content and atmospheric pressure (Lloyd, 1991; Thomas *et al.*, 1999; Wang
112 *et al.*, 2017a). But as λ is not directly measurable, Eqn 1 lacks useful predictive power over time
113 scales longer than the diurnal cycle. In consequence, it has not been adopted in models, which
114 instead typically constrain g_s using the empirical observations of a strong positive correlation
115 with A (Wong *et al.*, 1979) and a strong negative response to D .

116 Recent studies have attempted to combine theoretical and empirical approaches. Medlyn *et al.*
117 (2011) presented a unified model of optimal stomatal conductance (g_s^*):

$$g_s^* \approx g_0 + 1.6 \left(1 + \frac{g_1}{\sqrt{D}}\right) \frac{A}{C_a} \quad (2)$$

118 where g_0 and g_1 are parameters that can be estimated using gas-exchange measurements with
119 varying D . Eqn 2 is an approximate solution of Eqn 1 under electron-transport limitation of
120 photosynthesis (adopted on the grounds that guard cells controlling stomatal aperture have
121 markedly greater capacity for electron transport than for Rubisco C fixation) (Medlyn *et al.*,
122 2011). The model's slope parameter g_1 is proportional to the product of λ and the
123 photorespiratory CO_2 compensation point (Γ^*). The value of g_1 was found empirically to vary
124 among PFTs and to increase with growth temperature in line with the increase in Γ^* (Medlyn *et*
125 *al.*, 2011; Lin *et al.*, 2015). In a parallel result obtained by Katul *et al.* (2010) for Rubisco-limited
126 photosynthesis, optimal stomatal behaviour described by Eqn 1 yields a similar response to D
127 but a different response to C_a (see discussion by Medlyn *et al.*, 2013).

128 From a modelling perspective, however, the chief difficulty with Eqn 2 is that we have no *a priori*
129 derivation of the parameter values g_0 and g_1 . Several alternatives to the Cowan-Farquhar
130 hypothesis have been proposed, but because rates of A and E respond independently to so
131 many variables a fully mechanistic understanding of the costs of stomatal opening is still lacking
132 (Dewar *et al.*, 2018). Sperry *et al.* (2016) and Wolf *et al.* (2016) considered an optimization
133 criterion whereby an explicit carbon cost is assigned to the risk of hydraulic failure associated
134 with water transport. Wolf *et al.* (2016) further critiqued the Cowan-Farquhar criterion on
135 theoretical grounds, as its implicit water cost is based on the loss of water from a store shared by
136 all plants – ignoring competition between plants for the same water violates the criterion for an
137 evolutionarily stable strategy. Prentice *et al.* (2014) introduced an optimization criterion that
138 assigns explicit carbon costs to the maintenance capacities for both water transport and
139 carboxylation, and assumes that plants minimize the fraction of assimilation assigned to the sum
140 of these costs. This criterion leads to an optimal value of χ (χ_o):

$$\chi_o \approx \frac{\xi}{(\xi + \sqrt{D})} \text{ with } \xi = \sqrt{\left(\frac{bK}{1.6a}\right)} \quad (3)$$

141 where K is the effective Michaelis-Menten coefficient of Rubisco (i.e. reflecting the enzyme's
142 twin affinities for CO_2 and O_2), and a and b are empirical cost factors for water transport and
143 carboxylation respectively. Mathematically, equation (3) has a similar form to Eqn 2. If g_0 in

144 Eqn 2 is ignored, re-arrangement yields the same optimal value of χ as Eqn 3 with $\xi = g_1$. But
 145 importantly ξ (again a parameter related to the water cost of carbon gain) now depends on K
 146 rather than Γ^* , and the term λ has been replaced by the ratio b/a , which can be estimated
 147 empirically from instantaneous gas exchange measurements (as in Lin *et al.*, 2015) or stable
 148 carbon isotope ratios ($\delta^{13}\text{C}$). χ_o increases with T because of the temperature-dependency of K
 149 (Bernacchi *et al.*, 2001) and because the viscosity of water has an inverse relationship with T , so
 150 that the cost factor for water transport (a) decreases with T (Prentice *et al.*, 2014). Wang *et al.*
 151 (2017b) showed that χ_o is a function of growth D , T and altitude (z) (Wang *et al.*, 2017a) and can
 152 be represented in a linearized form around $T = 25^\circ\text{C}$ as follows:

$$\ln[\chi_o/(1 - \chi_o)] = 1.189 + 0.0545 (T - 25) - 0.5 \ln D - 0.0815 z \quad (4)$$

153 in which the constant 1.189 (dependent on b/a) was estimated from a large global set of $\delta^{13}\text{C}$
 154 measurements, while the predicted coefficients of $(T - 25)$, $\ln D$ and z in equation (4) lay within
 155 confidence intervals estimated empirically from the $\delta^{13}\text{C}$ dataset. Under standard growth
 156 conditions ($T = 25^\circ\text{C}$, $D = 1 \text{ kPa}$, $z = 0 \text{ m}$), Eqn 4 yields $\chi_o = 0.77$ and this value decreases with
 157 D , increases with T , and decreases with z .

158 Eqn 4 provides testable predictions of χ , and therefore of C_i . But to predict A using the
 159 standard biochemical model of C_3 photosynthesis (Farquhar *et al.*, 1980; the FvCB model) as
 160 generally employed in ESMs, we also need to know the carboxylation capacity (V_{cmax}) and the
 161 rate of electron transport (J). Here we introduce another long-standing optimality hypothesis,
 162 the co-limitation or co-ordination hypothesis (e.g. Chen *et al.*, 1993; Maire *et al.*, 2012), which
 163 states that the two limiting rates of photosynthesis tend toward equality under average daytime
 164 conditions. By equating these two co-limiting photosynthetic rates and making the simplifying
 165 assumption that J is proportional to photosynthetically active radiation (PAR) absorbed by the
 166 leaf (I_L), Dong *et al.* (2017) derived an approximation of optimal V_{cmax} :

$$V_{\text{cmax}} \approx \varphi_0 I_L \frac{(C_i + K)}{(C_i + 2\Gamma^*)} \quad (5)$$

167 where φ_0 is the intrinsic quantum efficiency of photosynthesis, here taken to equal 0.085
 168 (Collatz *et al.*, 1991). The description of the leaf light term (I_L) is given in Methods (below).
 169 Pursuing the co-ordination hypothesis further, the ratio of the rate of maximal electron transport
 170 (J_{max}) to V_{cmax} also points to investment strategies within the leaf (Poorter & Evans, 1998) that

171 may in turn be influenced by the environmental conditions experienced by the plant. Further
172 derivations in Wang *et al.* (2017b) yield an optimal ratio:

$$\frac{J_{max}}{V_{cmax}} = \frac{4}{(C_i + K)} \cdot \sqrt[3]{\frac{(C_i - \Gamma^*) \cdot (C_i + 2\Gamma^*)^2}{0.41}} \quad (6)$$

173 where the constant 0.41 was estimated by assuming that typical values for the ratios J_{max}/V_{cmax}
174 (1.88) and χ (0.8) correspond with standard growth conditions (see references in Wang *et al.*).

175 In the current study, by employing independent methods for estimating the leaf's C_i , we set out
176 to test the validity of these foregoing predictions. In seeking to validate the models explicitly
177 related to stomatal behaviour (parameters g_i and χ), we focus on the estimates derived using
178 time-averaged data ($\delta^{13}C$ rather than instantaneous gas exchange) as the more useful test for
179 dynamic vegetation models on similar grounds to those proposed above, arguing that a constant
180 λ is infeasible. We present a unique Australian dataset that combines leaf gas-exchange
181 measurements with concurrent, site-based meteorological data and foliar carbon isotope analysis.
182 The dataset comprises 50 plant species growing at seven environmentally contrasting sites, with
183 repeat measurements taken in contrasting seasons. We used these data to address the following
184 questions:

- 185 1. How well do instantaneous (gas-exchange) and time-averaged (isotope) measures of
186 optimal stomatal behaviour agree?
- 187 2. Do the predictions of the unified stomatal model represented by Eqn 2 help to reveal
188 spatial and seasonal patterns of leaf water use efficiency?
- 189 3. Can temporal variation in the integrated C_i/C_a ratio, as indexed by $\delta^{13}C$ measurements,
190 be predicted from growth temperature and vapour pressure deficit as indicated by Eqn 4?
- 191 4. How well do predictions of V_{cmax} , and the ratio J_{max}/V_{cmax} , derived under the co-
192 ordination hypothesis following Eqns 5 and 6 correlate to independent gas-exchange
193 measurements?

194

195

196 **Methods**

197 *Study sites and climate data*

198 Our seven study sites are a subset of TERN Ecosystem Processes 'SuperSites'
199 (supersites.tern.org.au) (Karan *et al.*, 2016). TERN Ecosystem Processes is a platform of the
200 Terrestrial Ecosystem Research Network (TERN), Australia's land–ecosystem observatory. Site
201 locations and key descriptors of dominant vegetation and soil type are presented in Table 1. The
202 sites were chosen to provide a wide range of environmental conditions (Supporting Information
203 Fig. S1). Each SuperSite is equipped with a flux tower (TERN OzFlux) that records a common
204 suite of meteorological data (Beringer *et al.*, 2016). Our initial visits pre-dated the installation of
205 the standard OzFlux system at three sites; in those instances we used the ANUClimate statistical
206 model (Hutchinson *et al.*, 2009) and data from the Australian Bureau of Meteorology's nearest
207 weather station (TERN eMAST). The proportion of down-welling shortwave radiation (Fsd,
208 OzFlux) deemed as PAR was 0.47 (Britton & Dodd, 1976).

209 Apart from Alice Mulga, each site was visited twice. The timing of the visits was designed,
210 within logistical constraints, to provide the widest possible seasonal contrast. The prevailing
211 climate conditions leading up to each campaign are provided in Table 2. In testing for
212 environmental dependencies of leaf traits, we considered a range of time-periods. In a
213 glasshouse experiment on different *Nothofagus* species, Read & Farquhar (1991) found that for
214 Australasian species the environmental parameter of the collection site most closely correlated
215 with genetically determined C isotope discrimination (Δ) was the precipitation from the period
216 December to March. Our study of leaf traits included seasonal contrasts (i.e. sub-annual) and so
217 an intermediate (quarterly) time step has been adopted to represent average environmental
218 conditions. Model iterations (e.g. Eqns 4 and 5) were run also adopting weekly and monthly
219 time-steps, but offered neither improved explanatory power nor qualitative changes to the results
220 (data not shown).

221 *Leaf gas exchange*

222 The plants measured (50 species, Supporting Information Table S1), were selected to include
223 locally dominant species of trees and shrubs. This is a subset (combined here with C isotope
224 analysis) of a larger dataset presented in Bloomfield *et al.* (2018). At each visit, we chose young,
225 fully developed leaves from two sun-exposed branches. Leaf gas exchange measurements were
226 concentrated in the morning and performed using portable photosynthesis systems (Li-Cor 6400,
227 Li-Cor, Lincoln, NE, USA), using the 6 cm² chamber fitted with the red-blue light source (Li-

228 Cor 6400-02B LED). Upper canopy branches were excised using forestry shears on telescopic
229 poles and the branches immediately placed in a bucket and the branches recut under water to re-
230 establish the xylem water column (Domingues *et al.*, 2010). Performing gas-exchange
231 measurements on excised branches can affect subsequent calculations where stomatal
232 conductance is heavily depressed. Our initial data exploration excluded such outliers (36 rows for
233 individual leaves, 3.2% of the dataset).

234 For each leaf, approximately light-saturated ($1500 \mu\text{mol photons m}^{-2} \text{s}^{-1}$) measures of net
235 photosynthesis were taken at two CO_2 concentrations: $400 \mu\text{mol mol}^{-1}$ (ppm) (A_{400}) and 1500
236 $\mu\text{mol mol}^{-1}$ (A_{1500}). The leaf was next wrapped in aluminium foil for 30 minutes before
237 respiration in darkness (R_{dark}) was measured, still at 400 ppm CO_2 . Air flow was held constant
238 and a constant chamber block temperature (T_{Block}) was adopted for all measurements at a given
239 site and season, set marginally ($\approx 1^\circ\text{C}$) higher than expected morning air temperatures to counter
240 the effect of transpirational cooling and ensure leaf and ambient air temperatures were similar.
241 T_{Block} settings ranged from 10°C for the winter visit to Warra to 32°C for the summer visit to
242 Calperum, reflecting the wide range of air temperatures experienced across the sites and seasons.
243 The Li-Cor 6400 system can maintain T_{Block} values $\pm 6^\circ\text{C}$ away from ambient and this operating
244 constraint precluded measuring gas exchange at a common temperature across all sites and
245 seasons. With a constant flow rate, chamber humidity conditions varied and mean D within the
246 chamber ranged from 0.55 kPa for the winter visit to Warra to 4.02 kPa for the summer visit to
247 Calperum.

248 In the absence of $A-C_i$ response curves, we estimated V_{cmax} based on our light-saturated A_{400}
249 values using the ‘one-point method’ described and validated by de Kauwe *et al.* (2016).
250 Mitochondrial respiration in the light (R_{day}) was assumed to be equivalent to R_{dark} (this simplifying
251 assumption has only a very minor effect on the estimation of V_{cmax}). The rate of electron
252 transport at saturating $[\text{CO}_2]$ of 1500 ppm (J_{max}) was estimated according to the equation:

$$J_{\text{max}} = (A_{1500} + R_{\text{dark}}) \cdot \frac{(4C_i + 8\Gamma^*)}{(C_i - \Gamma^*)} \quad (7)$$

253 Our estimates of V_{cmax} and J_{max} reflect temperature dependencies applied to K and Γ^* according
254 to Bernacchi *et al.* (2001). To allow comparisons of metabolic rates across sites and seasons, we
255 calculated rates at both a standard temperature (25°C) and at the average daytime air temperature
256 for the 90-day period leading up to the end of each field campaign (Table 2). From the rate
257 derived at a given measurement temperature, normalised V_{cmax} was calculated by applying the

258 Arrhenius function (Medlyn *et al.*, 2002) assuming an activation energy of 64.8 kJ mol⁻¹ (Badger
259 & Collatz, 1977).

260 ***Photosynthetic model extensions***

261 When modeling V_{cmax} (Eqn 5), we adopted a simple leaf light term (I_L):

$$I_L \approx I_0 \frac{(1 - e^{-k \text{LAI}})}{\text{LAI}} \quad (8)$$

262 that estimates mean canopy light availability based on leaf area index (LAI) in the plot, where I_0
263 is incident PAR at the top of the canopy and k is an extinction coefficient. We have assumed k
264 = 0.5 since even distal leaves on sun-exposed branches usually experience only a fraction of full
265 PAR incident at the horizontal plane (e.g. Kitajima *et al.*, 2005). As LAI approaches zero, the I_L
266 term's upper limit is $I_0/2$, but becomes a diminishing fraction of I_0 as canopy density increases.
267 Even though our sampling protocol called for sun-exposed leaves, we consider that adopting I_0
268 would overestimate PAR for many mid-canopy or understory species.

269 The FvCB model and its subsequent implementations make the simplifying assumption for C_3
270 plants that CO₂ conductance through the leaf's mesophyll (g_m) is infinite such that the partial
271 pressure at the site of carboxylation in the chloroplast (C_c) is equal to C_i . It is known, however,
272 that internal resistances can lower g_m and Evans & von Caemmerer (1996) reported that for
273 many species C_c may be around 30% lower than C_i when leaves are actively photosynthesising in
274 high light. Variations in g_m are potentially important in a study such as ours that includes
275 measurements of plants growing at environmentally contrasting sites and at unfavourable
276 seasons. Following the approach in Ayub *et al.* (2011) we estimated g_m based on the strong
277 empirical correlation with A_{sat} presented in Evans & von Caemmerer (1996), but adopting a
278 drawdown (sub-stomatal airspace to chloroplasts: A/g_m) of 50 $\mu\text{mol mol}^{-1}$ suggested as a lower
279 bound in a subsequent meta-analysis by Niinemets *et al.* (2009). C_c was then derived according
280 to:

$$C_c = C_i - \frac{A_{\text{sat}}}{g_m} = C_i - 50 \quad (9)$$

281 Photosynthetic parameters inferred from gas exchange measurements ('one point' estimates of
282 V_{cmax} ; J_{max} according to Eqn 7) are presented in the analysis that follows after substituting C_c for
283 C_i and adopting lowered Michaelis constants for Rubisco's carboxylation and oxygenation
284 reactions after von Caemmerer *et al.* (1994). Dropping the assumption of infinite g_m , Wang *et al.*

285 (2017b) derived a modified equation for χ_c with unchanged environmental dependencies, but
286 producing a lower constant of 1.097 (replacing the 1.189 intercept term in Eqn 4) reflecting the
287 anticipated drawdown from C_i to C_c i.e. $\chi_c < \chi_o$.

288 ***Estimating χ from stable carbon isotope discrimination***

289 As an alternative to instantaneous methods, χ can be inferred from the leaf's stable isotope
290 composition. In their uptake of atmospheric CO_2 , plants discriminate against the stable ^{13}C
291 isotope relative to the lighter, more abundant ^{12}C isotope and the level of discrimination is
292 influenced by environmental conditions (Farquhar *et al.*, 1989). Isotopic discrimination (Δ) can
293 thus be expressed as the $^{13}\text{C}/^{12}\text{C}$ isotope ratio ($\delta^{13}\text{C}$) of leaf tissue relative to that of source air.
294 We here employ the simple model developed by Farquhar *et al.* (1982) whereby discrimination is
295 given by:

$$\Delta (\text{‰}) = \frac{\delta^{13}\text{C}_a - \delta^{13}\text{C}_l}{1 + \delta^{13}\text{C}_l/1000} \quad (10)$$

296 where $\delta^{13}\text{C}_a$ is the $\delta^{13}\text{C}$ value of CO_2 in air (assumed to be -8‰) and $\delta^{13}\text{C}_l$ is that of the leaf. It is
297 then possible to derive time-averaged values of χ (weighted for photosynthetically active periods)
298 based on the relationship with Δ according to:

$$\chi = \frac{(\Delta - a')}{(b' - a')} \quad (11)$$

299 where a' is the isotope fractionation arising from diffusion through the stomata and b' the net
300 fractionation caused by carboxylation, with standard values for C_3 plants of 4.4‰ and 27‰
301 respectively. Eqn 11 is an approximation and disregards a number of additional dependencies of
302 $\delta^{13}\text{C}$ (e.g. Ubierna & Farquhar, 2014), but works well in practice because the generally adopted
303 value of b' implicitly includes these effects (Cernusak *et al.*, 2013). An extended formulation that
304 included explicit fractionation terms for mesophyll transfer and photorespiration was presented
305 by Seibt *et al.* (2008). We here extend their 'classical' model to include a term presented by
306 Farquhar & Cernusak (2012) that corrects the isotopic estimates of C_i for the effects of
307 transpiration on the diffusive flux of CO_2 already captured in the LiCor formulation – the
308 ternary factor (so called to describe the collision of molecules of CO_2 , water vapour and air).
309 Solving for χ provides a revised ratio:

$$\chi = \frac{\Delta(1-t) - a' + (b - a_m)(1+t)\frac{A}{g_m C_a} + (1+t)f\frac{\Gamma^*}{C_a}}{(b + tb - a')} \quad (12)$$

$$\text{where } t = \frac{E}{2\frac{g_s}{1.6}}$$

310 is the ternary correction factor, b is the mean fractionation during carboxylation by Rubisco
 311 (30‰), a_m is fractionation during transfer through the mesophyll (1.1‰) and f is fractionation
 312 during photorespiration (8‰). Hereafter we distinguish between these two discrimination
 313 models by referring to the simple (Eqn 11) and neo-classical (Eqn 12) formulations.

314 $\delta^{13}\text{C}$ methods, conducted on leaves sampled from the same branch as gas exchange
 315 measurements, are as described in Rumman *et al.* (2018). Briefly, after drying each leaf sample
 316 was finely ground until homogenous with a ball mill (Retsch MM300, Verder Group,
 317 Netherlands). Between one and two milligrams of ground material were sub-sampled in tin
 318 capsules for analysis of $\delta^{13}\text{C}$, giving three representative values per tree. All $\delta^{13}\text{C}$ analyses were
 319 performed using a Picarro G2121-i Analyser (Picarro, Santa Clara, CA, USA) for isotopic CO_2 .
 320 Atropine and acetanilide were used as laboratory standard references and results were normalised
 321 with the international standards sucrose (IAEA-CH-6, $\delta^{13}\text{C}_{\text{VPDB}} = -10.45$ ‰), cellulose (IAEA-
 322 CH-3, $\delta^{13}\text{C}_{\text{VPDB}} = -24.72$ ‰) and graphite (USGS24, $\delta^{13}\text{C}_{\text{VPDB}} = -16.05$ ‰). Standard deviation
 323 of the residuals between the IAEA standards and the calculated values of $\delta^{13}\text{C}$ based on best
 324 linear fit was ≈ 0.5 ‰.

325 **Estimating g_1**

326 The value of $g_{1\text{-leaf}}$ was estimated by fitting the unified stomatal model (Eqn 2) to our leaf gas
 327 exchange data using the fitBB function in the *plantecophys* R package (Duursma, 2015); gas
 328 exchange data for input to the model were restricted here to the A_{400} measurements described
 329 above. Values of isotope-derived g_1 were estimated following Medlyn *et al.* (2011) using χ values
 330 obtained from Eqn 12 according to

$$g_{1\text{-isotope}} = \frac{(\chi \sqrt{D})}{(1 - \chi)} \quad (13)$$

331 where D relates to time-averaged daylight site values obtained from the OzFlux network.

332 **Data analysis**

333 All statistical analyses were conducted using the open-sourced statistical environment 'R' (R
334 Development Core Team, 2017). Leaf trait and parameter observations are presented as species
335 averages for a given site and season; sample size and standard deviations are given in Supporting
336 Information (Table S2). The level of agreement between two independent measures of a trait, or
337 between model predictions and independently measured trait values, was tested using Pearson
338 correlations. Linear mixed effects models, from the *nlme* package (Pinheiro *et al.*, 2018), were
339 used to test for differences between sites and seasons for the g_1 parameter and for relationships
340 between isotope-derived χ and the environmental variables T and D ; a common random effect
341 allowed different intercepts for each plant family. Differences between sites were assigned using
342 Tukey's honest significant differences (HSD). Comparisons between nested models were
343 assessed using log likelihood ratios (L). Partial residual plots showing the relationship between
344 independent parameter estimates and model explanatory variables were represented using the
345 *visreg* R package (Breheny & Burchett, 2017).

346

347

348 **Results**

349 ***Stomatal behaviour***

350 We found a positive, but weak correlation ($r = 0.290$) between our gas exchange- and isotope-
351 derived versions of the g_1 slope parameter (Fig. 1) although for tropical forest species several of
352 the isotope-derived values were unrealistically high ($>20 \text{ kPa}^{0.5}$; corresponding to C_i/C_a ratios
353 approaching unity). Instantaneous estimates of the C_i/C_a ratio within the Li-Cor chamber were
354 consistently lower than the isotope-derived ratios, but the effect on $g_{1-\text{leaf}}$ was partially offset by
355 higher D within the chamber at the time of measurement relative to the daytime mean D of the
356 preceding quarter (Supporting Information Fig. S2). We found no seasonal differences in the
357 species-averaged $g_{1-\text{isotope}}$ estimates (even when models were run for a non-tropical subset of the
358 sites); stepwise deletions from a fully interactive Site x Season fixed term (Site + Season +
359 Site:Season) confirmed that the parsimonious Site-only model performed just as well ($L = 11.82$,
360 $df = 6$, $p = 0.066$). Wide variation among co-occurring species meant there were few differences
361 between the study sites, but $g_{1-\text{isotope}}$ values for Warra were lower than for the tropical forests (Fig.
362 2).

363 The expected influence of environmental conditions on stomatal behaviour was only partially
364 borne out. The $g_{L_isotope}$ parameter showed a weak positive relationship with time-averaged air
365 temperatures ($R^2 = 0.10$ Supporting Information Fig. S3), but was unrelated to available indices
366 of water availability: precipitation, moisture index or top-layer soil water fraction (e.g. for a
367 multiple regression style model, dropping the soil water term to leave temperature as the sole
368 explanatory variable made little difference to the model's performance: $L = 1.29$, $df = 1$, $p =$
369 0.257). Despite the limited agreement between our two versions of the g_1 parameter similar
370 patterns were observed for g_{1_leaf} (gas-exchange): a Site-only model was again preferred ($L = 9.07$,
371 $df = 6$, $p = 0.169$); inter-site differences here suggested a progression of arid < temperate <
372 tropical forests (Fig. 2 inset); but relationships with environmental variables were tentative at
373 best: T ($R^2 = 0.11$).

374 *Leaf Ci:Ca ratio*

375 Estimates of χ modelled as a function of environmental parameters T and D using Eqn 4 (for
376 this Australian dataset, we have ignored the limited site differences in altitude) showed strong
377 correlation ($r = 0.702$) with corresponding ratios derived from ^{13}C discrimination (Eqn 12) (Fig.
378 3); the correlation was stronger still when ratios for co-occurring species were expressed as a site
379 mean ($r = 0.930$). The modelled ratios were, however, consistently lower and we next
380 considered how well Eqn 4 described our Δ -derived ratios. The consistent offset from the
381 published model meant that our isotopic data yielded a markedly higher intercept at standard
382 conditions from that proposed by Wang *et al.* (2017b) - corresponding to $\chi=0.85$ (*versus* 0.75); the
383 slope of the T response did not differ significantly from the theoretical value, but the slope
384 coefficient for D was more steeply negative (Table 3). Our result of consistently higher Δ -
385 derived χ stems from the decision to adopt an expanded formula of ^{13}C fractionation; a
386 comparison of neo-classical *versus* simple $\chi_{isotope}$ confirms that, whilst very closely correlated, the
387 ratios were always higher under the extended formulation (Supporting Information Fig. S4b).

388 The fit of the Δ -derived ratios to the environmental data was improved when the two seasons
389 were treated separately. The independent effects of the model's two environmental parameters
390 suggested a seasonal shift in relative explanatory power with $T > D$ in the favourable season, but
391 that ranking reversed in the off-season linked to pronounced deficits at Calperum and GWW
392 (Table 3, Fig. 4).

393 **Photosynthetic parameters**

394 The maximal rate of Rubisco carboxylation (V_{cmax}), modelled using Eqn 5 adopting C_i estimates
395 derived from ^{13}C discrimination, showed fair agreement with estimates inferred from gas-
396 exchange measurements when expressed at the prevailing measurement temperatures ($r = 0.675$,
397 Fig. 5a). It is noticeable, however, that for a given site and season the rates inferred from gas
398 exchange were much more variable than the modelled values, reflecting the model's use of
399 environmental variables common to all co-occurring species.

400 The V_{cmax} rates shown in Fig. 5a reflect enzyme kinetics at the prevailing temperatures for each
401 site and season: average daytime temperatures (Table 2) for the modelled rates and leaf
402 temperatures close to a corresponding T_{Block} for the gas exchange measurements. When rates
403 were expressed normalised to a standard temperature of 25°C , the correlation between inferred
404 and modelled values proved weaker ($r = 0.348$, Fig. 5b). This disparity was caused, in part at
405 least, by the respective temperature differentials from the 25°C reference: leaf temperatures in
406 the Li-Cor chamber (inferred rates) always exceeded the daytime air temperatures (modelled
407 rates) (Methods). The divergence between inferred and modelled normalised rates became
408 more pronounced as the temperature differential increased (Supporting Information Fig. S5).
409 Once normalised for average canopy light conditions (I_L above), our 25°C standardised gas
410 exchange rates showed a negative relationship with growth temperature (Supporting Information
411 Fig. S6).

412 The modelled ratio $J_{\text{max}}/V_{\text{cmax}}$ (Eqn 6) similarly derived from independent isotope and
413 environmental variables provided a good fit to the gas exchange estimate (Fig. 6) although the
414 modelled ratios were generally higher – both ratios estimated at the prevailing temperature. For
415 C_3 species average $J_{\text{max}}/V_{\text{cmax}}$ ratios of around 2.0 are widely reported (e.g. Wullschleger, 1993),
416 but the model yielded many higher ratios (approaching 4.0 for the winter visit to Warra) and this
417 was related to cooler growth temperatures for certain sites and seasons – consistent with both
418 experimental evidence (Kattge & Knorr, 2007) and predictions (Wang *et al.*, 2017b). $J_{\text{max}}/V_{\text{cmax}}$
419 ratios inferred from gas exchange measurements confirmed a strong negative dependency on leaf
420 temperature, approaching unity as leaf temperatures neared 35°C (Supporting Information Fig.
421 S7).

422

423

424 Discussion

425 Theoretical optimality approaches permit prediction of photosynthetic parameters for species
426 and ecosystems for which few or no empirical data are available, but of course such hypotheses
427 require validation across as broad an environmental scale as possible. For diverse plant species
428 growing at contrasting sites across Australia we show that leaf physiological parameters key to
429 many Earth system models (ESMs) are predictable, to first order, using readily available
430 environmental variables. Collectively, our findings support the contention that there now exists
431 a unifying theoretical basis that ESMs can exploit to predict key leaf parameters across diverse
432 species and biomes (Wang *et al.*, 2017b).

433 *The unified stomatal model*

434 It is often assumed that estimates of the same parameter obtained using different methods
435 should match (e.g. Quentin *et al.*, 2015), but we found only weak agreement between our two
436 versions of g_1 . Although the underlying patterns of variation were similar for both bases, the
437 poor correlation prompts the question which is the more reliable or informative. Our gas-
438 exchange protocol was designed primarily to provide an instantaneous estimate of
439 photosynthetic capacity. It may be that for many of these species xylem performance was
440 compromised by the act of cutting the branches and hence water transport to the leaves was
441 disrupted causing a reduction in g_s (Santiago & Mulkey, 2003). By contrast, g_1 values derived
442 from the isotopic composition of leaf biomass reflect the growth environment as integrated over
443 the leaf's development. In addition, metabolic processes post photosynthetic fixation create
444 further fractionation steps that may cause Δ (hence χ and g_1) derived from leaf material to differ
445 from instantaneous estimates (Cernusak *et al.*, 2013).

446 A recent study by Medlyn *et al.* (2017) compared g_1 estimates derived from three different
447 datasets: gas exchange, leaf isotope (as here) and eddy covariance. They found that the three g_1
448 estimates were inconsistent, but in contrast to our results values for $g_{1\text{-isotope}}$ were generally lower
449 than $g_{1\text{-leaf}}$ for a given PFT. Our own study is heavily dominated by evergreen broad leaf trees (47
450 of 50 species, Supporting Information Table S1), a classification that straddles three of the PFTs
451 adopted in the Medlyn *et al.* study. For those three PFTs and shrubs (also represented here),
452 Medlyn *et al.* (2017) reported significant differences between $g_{1\text{-leaf}}$ and $g_{1\text{-isotope}}$ values only for the
453 savanna plants; the authors ascribe high variability in $g_{1\text{-leaf}}$ for savanna trees to marked seasonal
454 differences. Their meta-analysis, however, included measurements from savannas in the
455 seasonally dry tropics where monthly rainfall can vary by hundreds of mm (see for example Fig.
456 1 in Eamus *et al.*, 2000). The g_1 parameter is inversely proportional to the plant's intrinsic water

457 use efficiency (defined as A/g_s) and we expected to find inter-site differences that followed a
458 gradient of moisture availability. Instead differences emerged only between the temperate forest
459 at Warra and the moist tropical forests in FNQ. In a global study of 314 species at 56 sites, Lin
460 *et al.* (2015) found no difference in g_i among C_3 plants from temperate or tropical biomes, but
461 did find important differences among PFTs: tropical moist forest trees at 3.37, shrubs at 4.22
462 and evergreen savanna trees at 7.18. Compared to the Lin *et al.* study, our tropical forest results
463 are roughly double, whilst our g_i values for the arid sites, derived in part from evergreen savanna
464 trees, are much lower. Those lower g_i estimates reported here (Fig. 2) arguably conform better
465 to expectations of higher WUE in the dry conditions of those arid sites.

466 Stomatal behaviour must, in part, depend on environmental conditions, but we found no
467 evidence of a seasonal effect on values of g_i . Lack of a seasonal response might be expected for
468 moist tropical forests where the T and D differentials of wet *versus* dry seasons are much less
469 pronounced than the summer *versus* winter contrasts of mid-latitude sites (Table 2) (Buchmann *et*
470 *al.*, 1997). Surprisingly our results showed no relationship between g_i and indices of water
471 availability. In another Australian study of five evergreen eucalypt species growing at three sites
472 along a rainfall gradient in the Northern Territory, Eamus *et al.* (2000) found no seasonal
473 differences in transpiration rates despite marked seasonality in rainfall, aridity and leaf water
474 potential. The authors propose that for these savanna trees, water use is determined by the
475 limiting dry season conditions.

476 Our results showed poor consistency between the gas-exchange and isotope methods of
477 estimating optimal stomatal behaviour. We found that the g_i parameter of WUE was weakly
478 influenced by air temperature over broad scales, but locally was at once widely variable among
479 co-occurring species and conserved across changing seasons.

480 ***Predicting the ratio of leaf internal to ambient [CO_2]***

481 A large body of work from the 1980s onwards has shown that ^{13}C discrimination is a reliable
482 proxy for time-averaged χ (see Methods). Our results lend support to the model presented by
483 Wang *et al.* (2017b) and show that isotope-derived χ is closely correlated to the corresponding
484 value predicted from climate variables of growth T and D . In contrast to the original study,
485 however, we found that the isotope-derived ratios were consistently higher than those produced
486 by the model. Departing from the original study, we have applied an extended formulation of
487 ^{13}C fractionation (Eqn 12); the ternary and photorespiration corrections here are relatively minor
488 and act as partial offsets and so the principal disparity stems from our treatment of mesophyll
489 conductance. Current ESMs, and the standard implementation of the FvCB model, ignore the

490 limitation placed on photosynthesis by g_m . In this study we adopted a single empirical correlation
491 for estimating the drawdown from C_i to C_c (Eqn 9), but this is inevitably an oversimplification.
492 Despite difficulties in measurement, g_m is known to vary among species and in response to
493 environmental conditions (Flexas *et al.*, 2008; von Caemmerer & Evans, 2015). Importantly, ^{13}C
494 discrimination methods can be combined with gas exchange measurements to provide estimates
495 of g_m (Cernusak *et al.*, 2013). This is an important area for future research, both empirical and
496 theoretical, on the control of photosynthetic traits although debate remains as to the necessity of
497 incorporating g_m estimates into wider models (e.g. Bahar *et al.*, 2018).

498 Fitting the Wang *et al.* model to our data gave a coefficient of determination very similar to the
499 original study (R^2 of 0.41 *versus* 0.39), but the fit was improved if the two seasons were treated
500 separately (Table 3). One might expect seasonal contrasts to be captured by a stomatal model
501 that specifies T and D as explanatory variables, but we found that their relative importance
502 altered seasonally. In the favourable growing season, χ was more strongly influenced by
503 variation in daytime T , but in the unfavourable season the primacy shifted to D . Greater
504 influence for D under drier conditions is supported by a seasonal comparison of ecosystem
505 WUE at the Alice Mulga site, where the response to D was more pronounced in the dry (winter)
506 than the wet season linked to contrasting levels and variability of soil moisture content in the two
507 seasons (Eamus *et al.*, 2013).

508 Offset notwithstanding, the implication of our results is that we now have a means of estimating
509 C_i for modelling purposes. For this Australian study, where none of the sites sat higher than 750
510 m above sea level, we have ignored altitude effects on χ that are included in the original
511 formulation (Eqn 4). Further studies could usefully test the model along a gradient
512 encompassing a wider range of altitudes.

513 ***Photosynthetic parameters from first principles***

514 Whilst our predictions of key parameters provided generally good agreement with gas exchange
515 values, several discrepancies emerged. For a given site and season, the species-averaged V_{cmax}
516 rates inferred from our gas-exchange measurements showed wide variation. Such variability
517 amongst co-occurring species is to be expected (e.g. Bloomfield *et al.*, 2018) and likely reflects
518 factors such as canopy position, and therefore light environment, and VPD. Our model
519 performed less successfully when rates were normalised to the standard reference temperature of
520 25°C. The weaker correlation here may be partially explained by the underlying differentials in
521 temperature offset from the common reference (see above): leaf temperature *versus* daytime
522 averages. One effect of standardising the rates to a common reference temperature was to

523 narrow their range: inferred and modelled. Despite this contraction, the spectrum of site
524 averages indicated higher rates for the species growing in arid conditions. When our inferred
525 $V_{\text{cmax}25}$ rates were modelled as a function of the light, T and χ variables underpinning Eqn 5, the
526 residuals showed no discernible pattern with leaf mass per unit area (data not shown); suggesting
527 that variation in leaf morphology (at a coarse scale) provided little explanatory power beyond
528 these environmental variables.

529 What might be missing from our V_{cmax} model? In adopting a universal extinction coefficient
530 (recognising that not all leaves display perpendicular to the sun's rays) and a leaf PAR estimate
531 averaged for each plot (seasonal LAI values were adopted where available), our V_{cmax} model
532 assumed a uniform light environment for all leaves (Eqn 8). One of the major difficulties is that
533 we know very little about the light environments for these species. That quandary extends to our
534 gas-exchange measurements, which employed a single saturating light setting across the study,
535 and is likely one of the key reasons for the wide within-site variation in inferred V_{cmax} values. It
536 has, however, been demonstrated that within-canopy gradients of photosynthetic capacity and
537 light are not strictly proportional (e.g. Lloyd *et al.*, 2010) e.g. leaves in poor light appear to have
538 superfluous capacity thus breaching the optimality hypothesis. From our own results it is not
539 apparent that the model performed differently for those sites (Daintree and Cumberland Plain)
540 where we sampled tree branches from the gondola of a crane and where we might have expected
541 the light-averaging step to produce an underestimate. In considering why optimal canopies are
542 not observed in nature, Niinemets (2012) proposed a series of additional constraints or
543 mechanisms including leaf construction, acclimation and light-competition.

544 Even more challenging is the possibility that rates of photosynthesis are driven by the plant's
545 requirements for photosynthate (buffered by existing stores) rather than what is possible or
546 optimal. The demands - spanning respiration, reproduction, defence, exudates - represent a
547 complex network of plant processes intimately connected to the surrounding biotic and abiotic
548 environment and subject to stochastic fluctuation. The different timescales operating for our
549 two V_{cmax} estimates create the likelihood of quite distinct underlying phenology and sink
550 demand.

551 Changes or differences in growth temperature are expected to require adjustment in the plant's
552 photosynthetic apparatus, often manifested as variation in the $J_{\text{max}}/V_{\text{cmax}}$ ratio. Accordingly, we
553 found elevated ratios related to the cooler seasons and sites reflecting the ratio's negative
554 dependence on temperature (Fig. 6). Our gas exchange data suggested a strong negative linear
555 relationship with leaf temperature (Supporting Information Fig. S7), but with a steeper slope (-

556 0.08 *versus* -0.035 °C⁻¹) than that reported by Kattge & Knorr (2007). Our model (Eqn 6)
557 predicted J_{\max}/V_{cmax} ratios that were consistently higher than those inferred from the gas
558 exchange measurements and that coincides with the elevated χ_{isotope} values described above: the
559 modelled J_{\max}/V_{cmax} ratio increasing with C_i .

560 ***Concluding remarks***

561 We have used our unique dataset to test predictions of leaf photosynthetic traits across a range
562 of climates and plant species. Demonstrating the validity of these predictions has application for
563 the ESM community. Current ESMs are still hampered by uncertainty, much of which relates to
564 incomplete or inaccurate representations of the fundamental processes that govern carbon
565 fixation at the leaf, canopy and ecosystem levels. The path forward will involve sustained efforts
566 both to formulate explicit, unifying theories about the control of these processes and to test and
567 refine them with the help of field measurements. This research represents a step along that path.

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580

581

582 **Author contributions**

583 ICP, OKA and KJB designed the study. DE, DSE, MMB, PC, JC, MJL, CM, WSM, TW and
584 IJW co-ordinated site access, species identification and logistical support. KJB, JJGE, LSH,

585 HFT and LZ conducted fieldwork. BJE, MFH and MJL helped to compile the climate data. RR
586 conducted the isotope analysis and LAC proposed the extended discrimination model (Eqn 12).
587 BEM guided calculation of the g_1 parameter. KJB led the analysis of the data and wrote the
588 manuscript with input from all authors.

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788 **Supporting Information**

789

Fig. S1 Map and key climate indicators for each of the study sites.

Fig. S2 Instantaneous chamber conditions *versus* time-averaged equivalents.

Fig. S3 Parameter $g_{1_isotope}$ as a function of air temperature.

Fig. S4 Comparison of key parameters adopting ‘neo-classical’ and ‘simple’ fractionation.

Fig. S5 The effect of measurement temperature differentials on normalised rates of V_{cmax} .

Fig. S6 Temperature dependence of the V_{cmax} model Eqn 5.

Fig. S7 Temperature dependence of the $J_{\text{max}} : V_{\text{cmax}}$ ratio.

Table S1 Plant species by site.

Table S2 Leaf gas-exchange traits averaged by Site_Season_Species.

790

791

792 **Table 1** Site location and descriptors.

793

794 The sites are listed in order of increasing absolute latitude. An Australian map with the site locations indicated is included in Supporting Information Fig. S1. Leaf area
795 index is the ratio of projected leaf to ground surface area. Climate indices are annual long-term averages of interpolated data obtained from the Terrestrial Ecosystem
796 Research Network (TERN) eMAST facility. Moisture index is shown as the ratio of precipitation to potential evapotranspiration. Additional details for each site may be
797 found on the TERN SuperSite website (www.supersites.net.au). asl, above sea level; Leaf N, leaf nitrogen; Leaf P, leaf phosphorus.

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799 **Table 2** Climate conditions for each campaign (Site_Season) for the 90 d period ending on the final day of
800 fieldwork.

801

802 Air temperature (T_a), precipitation, net radiation (F_n), vapour pressure deficit (VPD), soil water fraction (S_{ws} ,
803 top layer), short-wave down-welling radiation (F_{sd}). Rainfall figures are totals, other values are means;
804 temperature, radiation and VPD figures relate to daylight hours only. In most cases the climate data have been
805 obtained from instruments centred on flux towers located at each site and managed under the OzFlux network
806 (www.ozflux.org.au). In three cases our fieldwork campaigns were made before the flux towers were fully
807 operational and in these instances we have used interpolated 24 h data obtained from ANUClimate
808 (www.emast.org.au/our-infrastructure/observations/anuclimate_data/). As we prepared the manuscript,
809 daily radiation data for the periods of interest were not available within ANUClimate and so those radiation data
810 have been obtained for the nearest automated weather station operated by the Australian Bureau of Meteorology
811 (www.bom.gov.au). The designation of Favourable and Unfavourable seasons is based on a local assessment of
812 growing conditions. na, not available.

813

814 **Table 3** Linear model summaries, following Wang *et al.* (2017b) and Eqn 4: in each case the response variable
815 is the logit-transformed isotope-derived ratio of leaf internal to ambient CO_2 partial pressure (χ).

816

817 The explanatory variables are the difference in the mean air temperature for the preceding quarter from $25^\circ C$
818 (ΔT , $^\circ C$) and the natural logarithm of the mean vapour pressure deficit for the preceding quarter ($\log_e D$, kPa).
819 The theoretical values are partial derivatives for each predictor, evaluated at 'standard' conditions ($T=25^\circ C$,
820 $D=1$ kPa). The t -test result is the probability of no difference between the fitted estimate and the theoretical
821 value (null hypothesis). Separate model iterations were run for the complete data-set and then for each of the
822 two seasons. na, not applicable.

823

824 **Fig. 1** Relationship between two estimates of the g_1 parameter (Eqn 2) derived from: instantaneous gas
825 exchange measurements (g_{1-leaf}) and leaf isotope discrimination ($g_{1-isotope}$). The points represent seasonal means
826 for individual plant species and sites are differentiated by colour. The dashed line represents the ideal 1 : 1 fit.
827 The x -axis has been restricted to exclude eight species (one Warra, five Robson Ck and two Daintree) with g_{1-}
828 $isotope$ greater than $20 \text{ kPa}^{0.5}$.

829

830 **Fig. 2** Boxplot showing the g_1 values estimated from leaf isotope discrimination grouped by site and including
831 both seasons. The y -axis has been restricted to exclude eight species (one Warra, five Robson Ck and two
832 Daintree) with $g_{1-isotope}$ greater than $20 \text{ kPa}^{0.5}$. The boxes indicate the interquartile range and median values.

833 Whiskers extend to the largest or smallest observation that fall within 1.5 times the box size; any observations
834 outside these values are shown as individual points. Boxes which share the same letter denote site means that
835 were not significantly different (Tukey's honest significant differences). The number of species included for
836 each site is indicated. The inset plot shows the equivalent analysis for the parameter derived from instantaneous
837 gas exchange measurements (g_{1-leaf} , notice that the scales of the axes differ).

838

839 **Fig. 3** Values of χ (the ratio of leaf internal to ambient CO₂ partial pressure) obtained from ¹³C isotope
840 discrimination, Eqn 12 plotted against corresponding predicted values of χ based on environmental variables of
841 temperature (T) and vapour pressure deficit (D) obtained from the OzFlux network, Eqn 4. Both estimates have
842 been modified to account for mesophyll conductance. The large coloured dots represent the site means
843 (combining all species and seasons) and the whiskers give the \pm SD. The smaller dots represent underlying
844 seasonal means for individual plant species (three to five plants per species). The dashed line represents the
845 ideal 1 : 1 fit. The Pearson correlation indicated relates to the underlying species' means (smaller dots); the
846 corresponding correlation based on site means (large dots) is 0.93.

847

848 **Fig. 4** Partial residual plots showing the logit-transformed ratio of leaf internal to external CO₂ partial pressure
849 (χ) derived from stable carbon isotope discrimination data as a function of time averaged environmental
850 variables: air temperature and vapour pressure deficit (D). Separate panels are shown for the contrasting visits:
851 (a, b) favourable season, (c, d) unfavourable season. Model coefficients and summaries appear in Table 3.

852

853 **Fig. 5** Goodness of fit plot for the carboxylation (V_{cmax}) model Eqn 5 using C_i data derived from carbon isotope
854 discrimination plotted against estimates inferred from independent gas exchange measurements. (a) Rates at the
855 prevailing measurement temperature; (b) the corresponding plot with rates standardised to a common reference
856 temperature of 25°C. The large coloured symbols represent the Site_Season means (combining all species) and
857 the whiskers give the \pm SD; the smaller dots represent underlying seasonal means for individual plant species
858 (three to five plants per species). The dashed line represents the ideal 1 : 1 fit. The Pearson correlations
859 indicated relate to the underlying species' means (smaller dots).

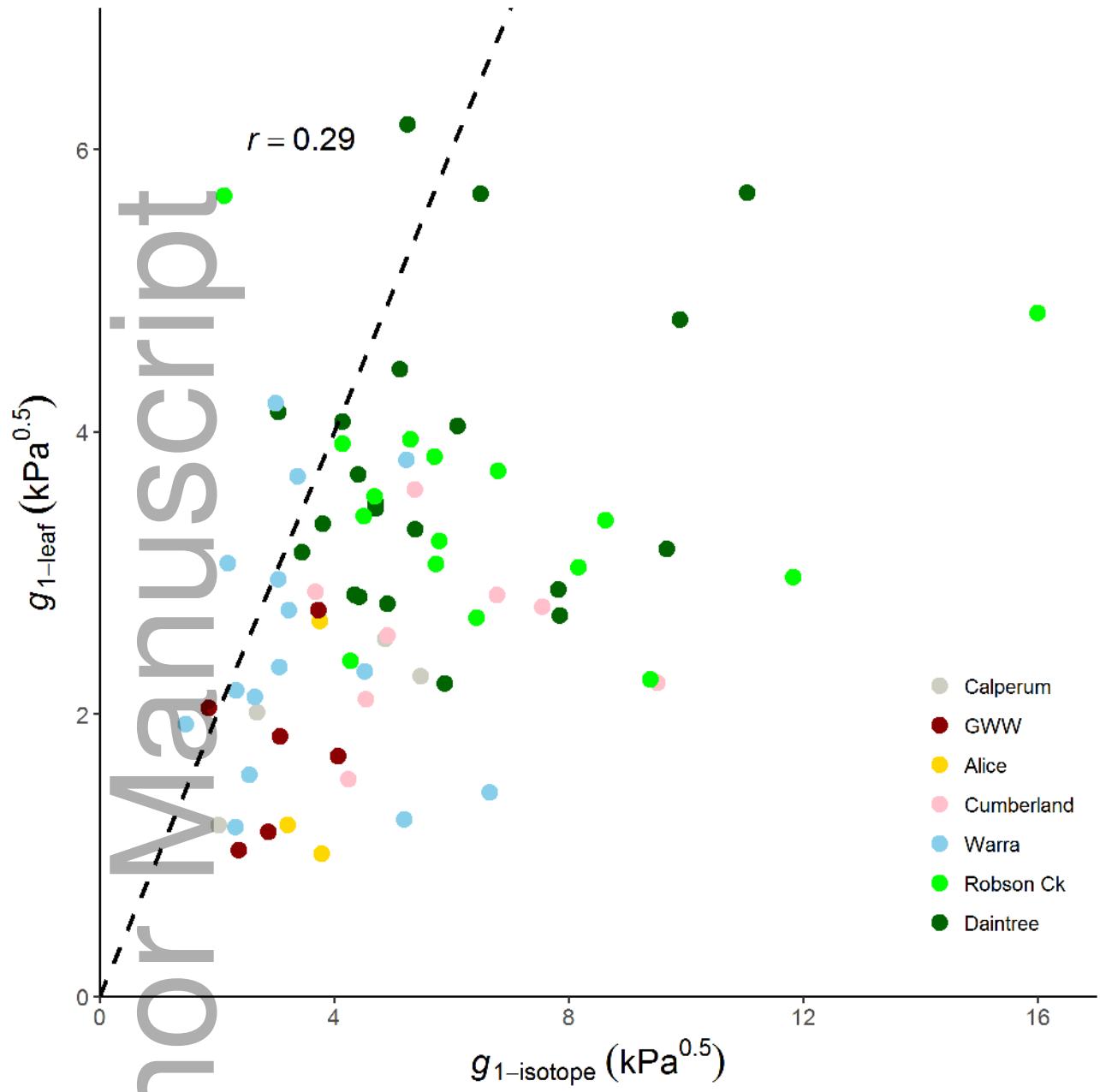
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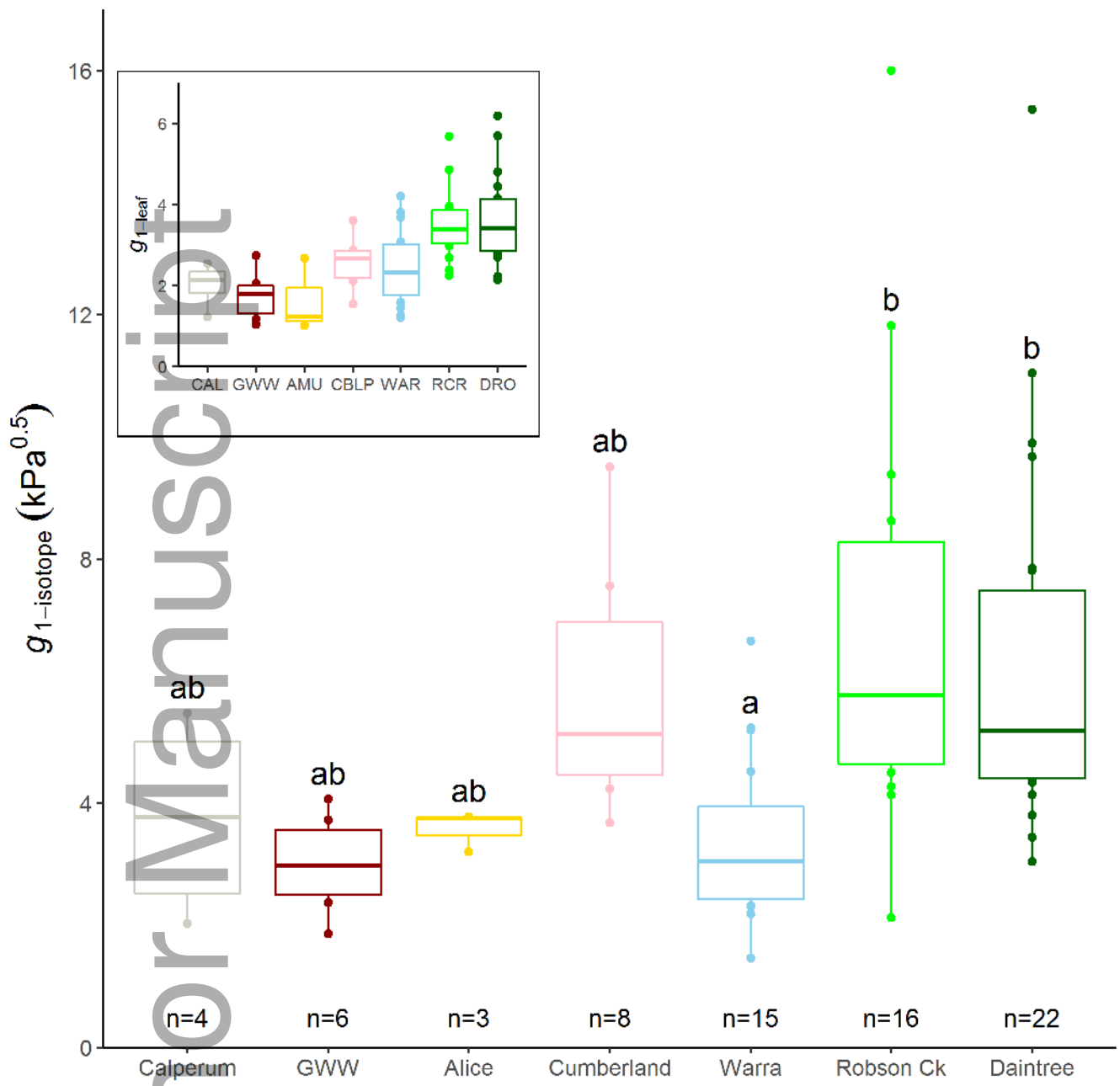
862 **Fig. 6** Goodness of fit plot for the modelled ratio of electron transport to carboxylation capacity ($J_{max} : V_{cmax}$)
863 from Eqn 6 using C_i data derived from carbon isotope discrimination plotted against independent gas exchange
864 estimates – both ratios expressed at the prevailing temperature. The large coloured symbols represent the
865 Site_Season means (combining all species) and the whiskers give the \pm SD. The smaller dots represent
866 underlying seasonal means for individual plant species (three to five plants per species). The dashed line

867 represents the ideal 1 : 1 fit. The Pearson correlation indicated relates to the underlying species' means (smaller
868 dots); the corresponding correlation based on Site_Season means (large symbols) is 0.86.

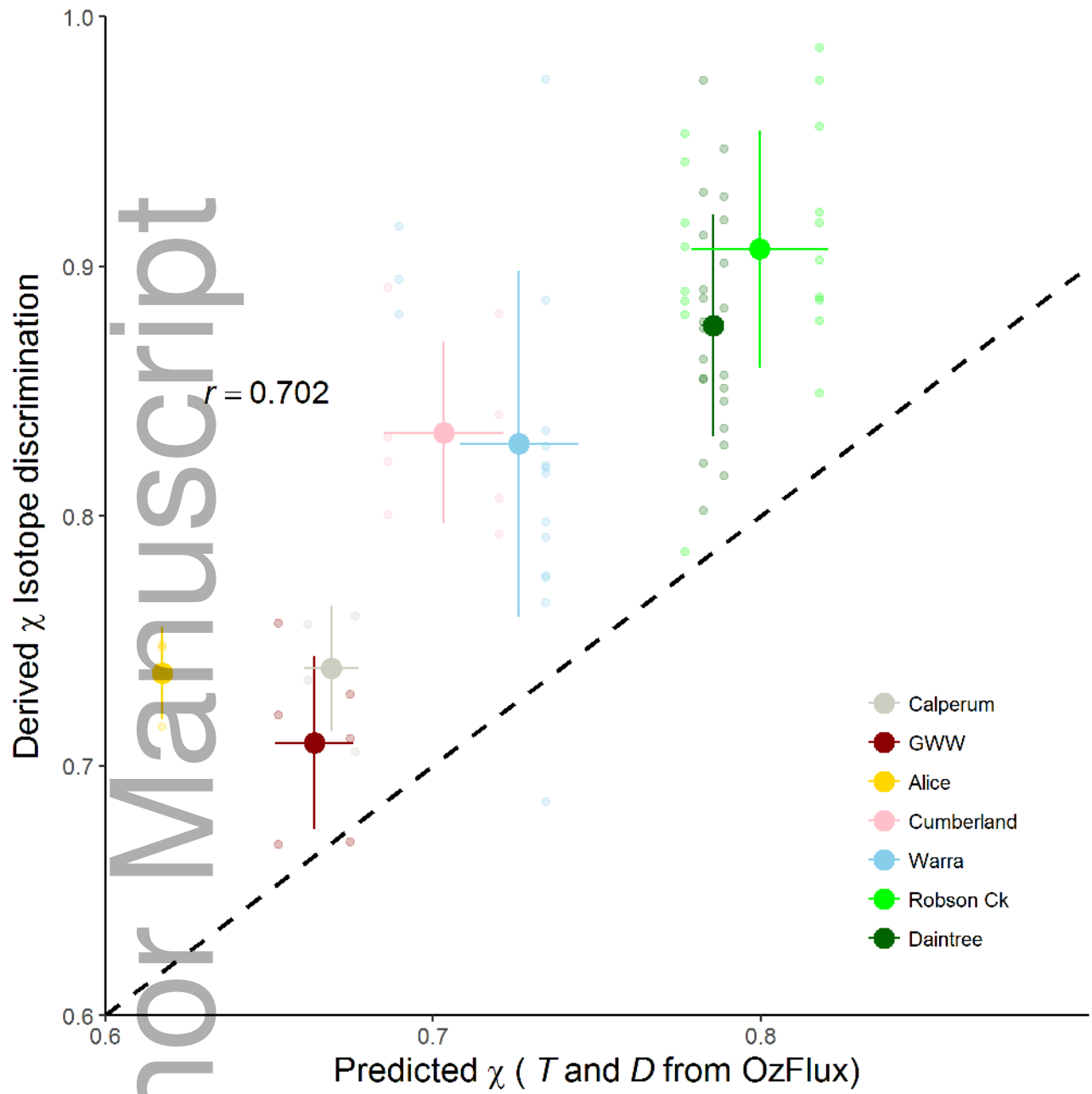
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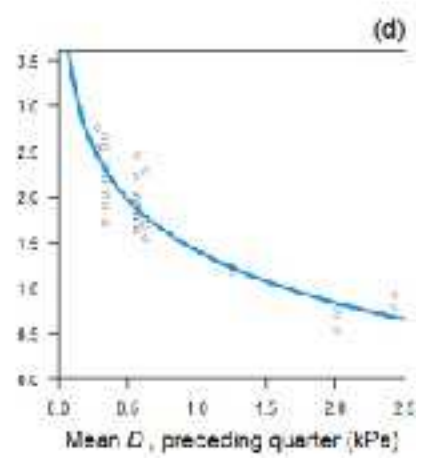
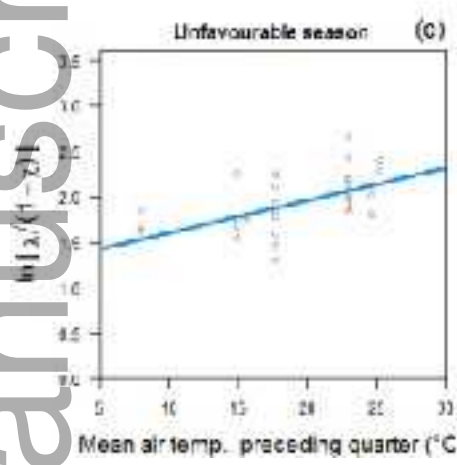
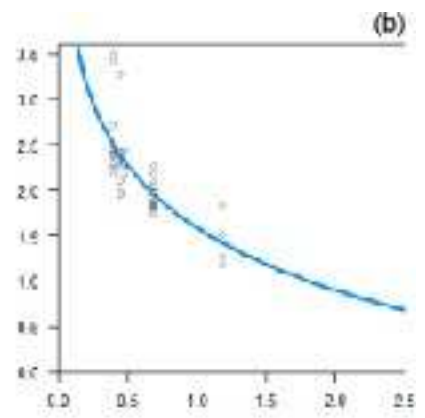
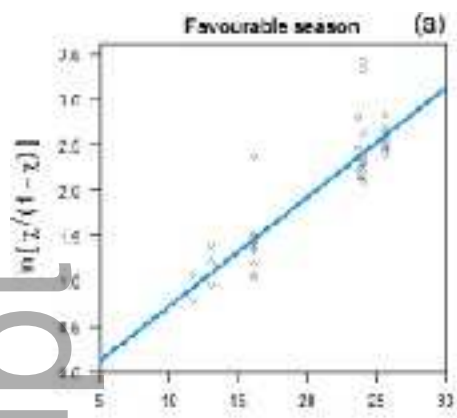
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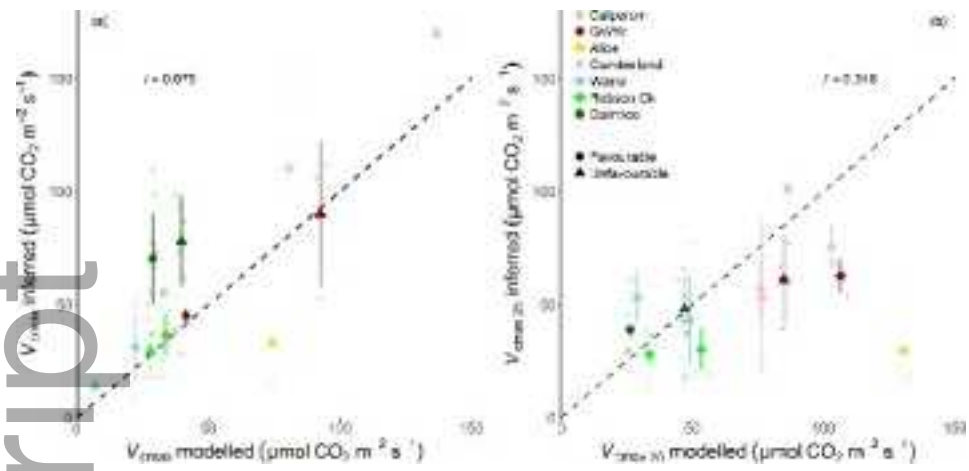
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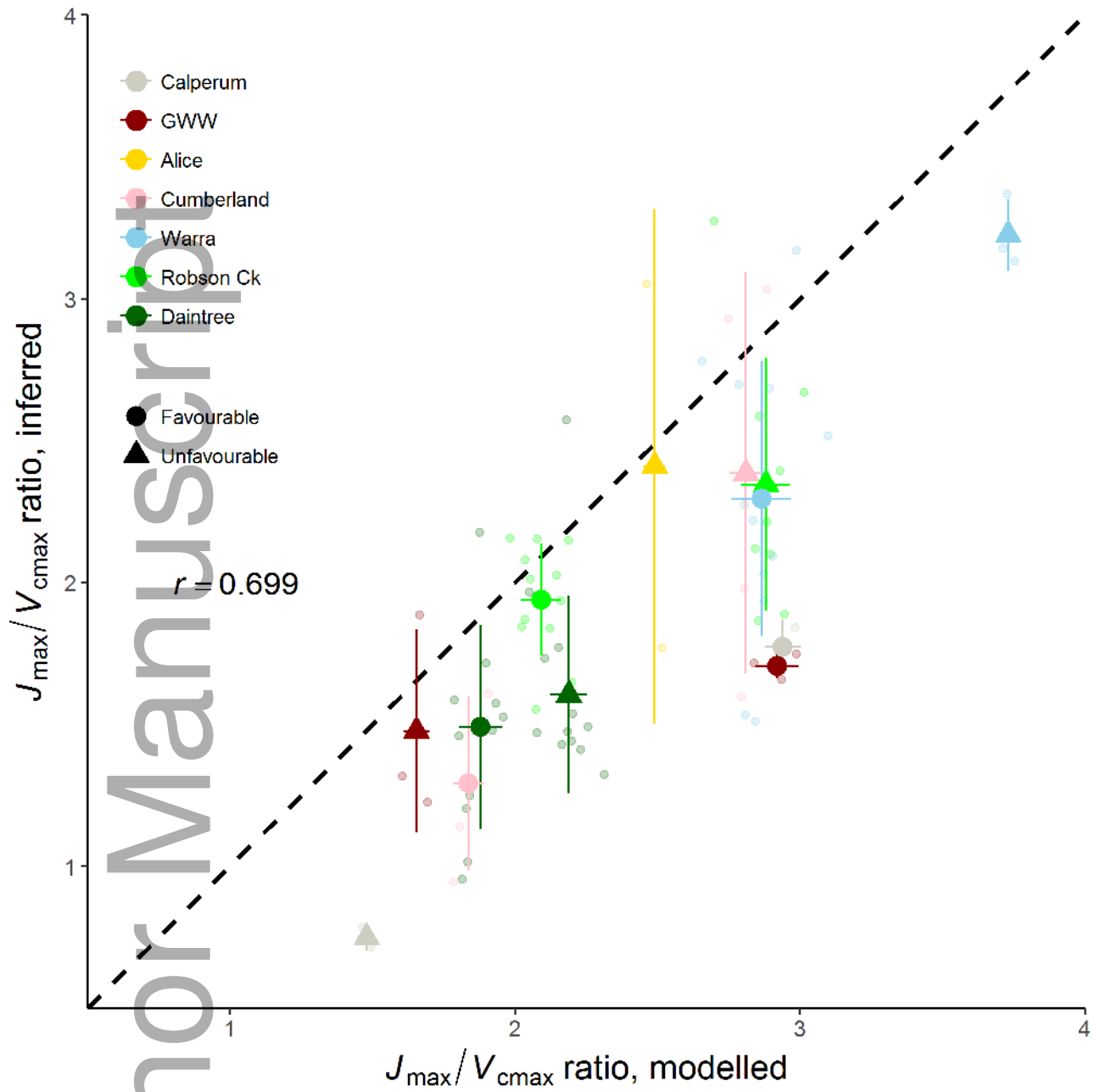
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